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## The Crystal Structure of Fairfieldite\*

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The mineral fairfieldite is triclinic, space group  $P\bar{1}$ ; the cell dimensions, as determined from crystals from Foote Mine, North Carolina, are:  $a=5.79$ ,  $b=6.57$ ,  $c=5.51$  Å,  $\alpha=102^\circ 16'$ ,  $\beta=108^\circ 40'$ ,  $\gamma=90^\circ 18'$ . The cell content is Ca<sub>2</sub>Mn(PO<sub>4</sub>)<sub>2</sub>·2H<sub>2</sub>O. The crystal structure was solved from a three-dimensional Patterson function and refined first by successive Fourier syntheses and then by least-squares to an  $R$  value 0.075. The coordination around the Mn and P ions is octahedral and tetrahedral respectively: the average Mn–O and P–O distances are 2.20 and 1.55 Å. The Ca atom has the coordination number 7+1 with seven Ca–O bonds in the range 2.34 to 2.50 Å and one bond 2.72 Å long. The structure consists of infinite chains of Mn octahedra and P tetrahedra running along the  $c$  axes. Calcium ions occupy the vacant spaces between the chains connecting them in a three-dimensional network. The crystal structure of the fairfieldite mineral group is closely related to that of the kröhnkite mineral group. In both structures there is the same chain motif with a strictly similar connection of tetrahedra and octahedra inside the chain. Differences occur in the reciprocal array of the chains, which are shifted in the  $c$  direction by different amounts in the two mineral groups.

### Introduction

The phosphate and arsenate minerals with the general formula A<sub>2</sub>B(XO<sub>4</sub>)<sub>2</sub>·2H<sub>2</sub>O, where A is Ca, B is Co, Fe, Mg, Mn, and X is As or P, may crystallize in the monoclinic or in the triclinic system. The monoclinic group comprises two minerals, brandtite and roselite, which are isostructural with the sulphate kröhnkite, CuNa<sub>2</sub>(SO<sub>4</sub>)<sub>2</sub>·2H<sub>2</sub>O. Members of the triclinic class are fairfieldite, neomesselite, collinsite and  $\beta$ -roselite.

The atomic arrangement in the monoclinic family was determined by Dahlman (1952) for kröhnkite and brandtite. Leone & Sgarlata (1954) redetermined the crystal structure of kröhnkite which was later refined by Rao (1961). In the present study the crystal structure of fairfieldite, which is a representative member of the triclinic group, was undertaken in an attempt to contribute to a better crystallochemical knowledge of phosphate minerals and to study its relationship with other compounds having a similar formula but belonging to the monoclinic system.

The mineral occurs as an accessory constituent in pegmatitic rocks, where it was first found by Brush

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& Dana (1879) at Branchville, Fairfield County (Connecticut). Wolfe (1940) gave the morphological elements of the crystals and the unit-cell dimensions, which were determined by the Weissenberg method.

On the basis of the chemical analysis, fairfieldite can be given the formula  $\text{Ca}_2(\text{Mn}, \text{Fe})(\text{PO}_4)_2 \cdot 2\text{H}_2\text{O}$ . There is a wide range of Mn/Fe atomic ratios: from 1.8:1 (Penfield, quoted in Brush & Dana) to 19.7:1 (Vassar, quoted in Landes, 1925).

For our investigation a small sample from Foote Mine, North Carolina, was used. The specimen consisted of white fairfieldite crystals with a pearly lustre, of short prismatic habit and associated with quartz.

### Experimental

Fairfieldite is not supplied by Foote Mine with a chemical analysis. Because of the small amount of material, a complete chemical analysis was impossible. However, before the X-ray study, the Mn/Fe ratio in the sample under investigation was determined by a microchemical colorimetric method, which gave a value of 49.8:1. This value is higher than those reported in the literature for fairfieldite from other localities.

A transparent single crystal of fairfieldite was ground into a nearly perfect sphere, 0.57 mm in radius. The sphere obtained was mounted on an integrating precession apparatus. Intensity data were collected with Zr-filtered Mo radiation for two sets of layers, taken within the mechanical limits of the apparatus:  $hkn$  ( $n=0$  to 2) and  $nkl$  ( $n=0$  to 3). The lattice constants determined from precession photographs are:

$$\begin{aligned} a &= 5.79 \pm 0.01 \text{ \AA} & \alpha &= 102^\circ 16' \pm 15' \\ b &= 6.57 \pm 0.01 & \beta &= 108^\circ 40' \pm 15' \\ c &= 5.51 \pm 0.01 & \gamma &= 90^\circ 18' \pm 15' \end{aligned}$$

in agreement with the values given by Wolfe (1940). The calculated density for one formula unit in the cell is  $3.095 \text{ g.cm}^{-3}$  which agrees with the value  $3.08 - 3.11 \text{ g.cm}^{-3}$  found in the literature.

The space group  $P\bar{1}$  was assumed at the first stage of the study and was confirmed by the crystal analysis.

Intensities were measured with a microdensitometer; Lorentz-polarization and absorption corrections ( $\mu R = 2.07$ ) were made using a program written by Catani & Zanazzi (1965) for the IBM 1620 computer. The  $F^2$  data were placed on the same relative scale by correlat-

ing the common reflexions on the films. 1089 non-equivalent diffraction effects were recorded, 185 of which were too weak to be observed.

### Structure determination and refinement

The crystal structure was solved by the three-dimensional Patterson function, and the successive Fourier syntheses method. The refinement was carried out by means of the least-squares full-matrix program of Busing & Levy, adapted for the IBM 7090 computer by Stewart (1964).<sup>\*</sup> Individual isotropic temperature factors were applied and the  $R$  index, defined as  $\sum ||F_o| - |F_c|| / \sum |F_o|$ , dropped from 0.25 to 0.075 in three cycles. Reflexions were given weights according to Hughes's (1941) scheme. Unobserved reflexions were excluded from calculations as well as some diffraction effects which seemed to be affected by extinction.

Anisotropic least-squares refinement did not yield an acceptable result because some thermal parameters shifted to give values without physical significance. A possible explanation for this fact may be an approximate absorption correction, the presence of secondary extinction or a slightly incorrect initial data rescaling. Isotropic temperature factors also assumed values lower than expected. Therefore, the refinement was stopped at this stage and no attempt was made to locate the hydrogen atoms.

The final atomic coordinates and individual isotropic thermal parameters are listed in Table 1 with their standard deviations. The observed and calculated structure factors are given in Table 2. The atomic scattering factors used in the calculations are those in *International Tables for X-ray Crystallography* (1962) for  $\text{Ca}^{2+}$ , P,  $\text{Mn}^{2+}$  and O. The scattering factor curve of  $\text{Mn}^{2+}$  was not corrected for the negligible Fe content.

### Discussion of the structure

The crystal structure of fairfieldite projected along the  $a$  axis is shown in Fig. 1. Bond lengths and angles with their standard deviations are listed in Tables 3 and 4.

<sup>\*</sup> Calculations were performed with the IBM 7090 computer of the Centro Nazionale Universitario di Calcolo Elettronico of Pisa University.

Table 1. Fractional atomic coordinates and isotropic thermal parameters

	$x$	$y$	$z$	$B$ ( $\text{\AA}^2$ )
Mn	0.0000	0.0000	0.0000	0.32 (3)
Ca	0.6020 (3)	0.2337 (2)	0.3332 (3)	0.22 (3)
P	0.2463 (3)	0.2325 (2)	0.6472 (3)	0.12 (3)
O(1)	0.0031 (11)	0.3264 (8)	0.6157 (11)	0.48 (7)
O(2)	0.4552 (11)	0.3995 (8)	0.6940 (10)	0.41 (7)
O(3)	0.2461 (11)	0.0699 (8)	0.3965 (10)	0.45 (7)
O(4)	0.3058 (10)	0.1177 (7)	0.8785 (10)	0.43 (7)
O(W)	-0.1543 (11)	0.3004 (8)	0.0663 (11)	0.58 (7)

Standard deviations are given in parentheses.



(Moore, 1967) it is 2.18 Å, in allactite (Moore, 1968a) it is 2.22 Å, in laueite (Moore, 1965) it is 2.12 Å and in clorophoenicite (Moore, 1968b) it is 2.21 Å.

The coordination number of the calcium ion can be assumed to be 7+1, each ion being surrounded by seven oxygen atoms at an average distance of 2.45 Å,

with values in the range 2.34–2.50 Å, while the next distance, Ca–O(3) (V), differs significantly from the others, with a value of 2.72 Å.

The PO<sub>4</sub> tetrahedra are nearly regular, with an average P–O bond length of 1.55 Å, a value in agreement with those found in metastrengite, 1.53 Å (Fanfani &

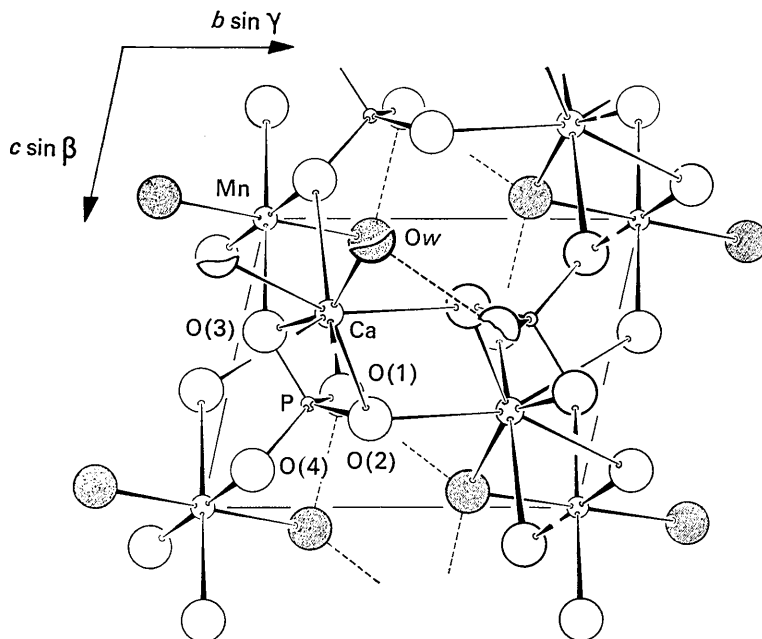


Fig. 1. Crystal structure of fairfieldite projected along the  $a$  axis. Shaded circles indicate water molecules,  $O(W)$ .

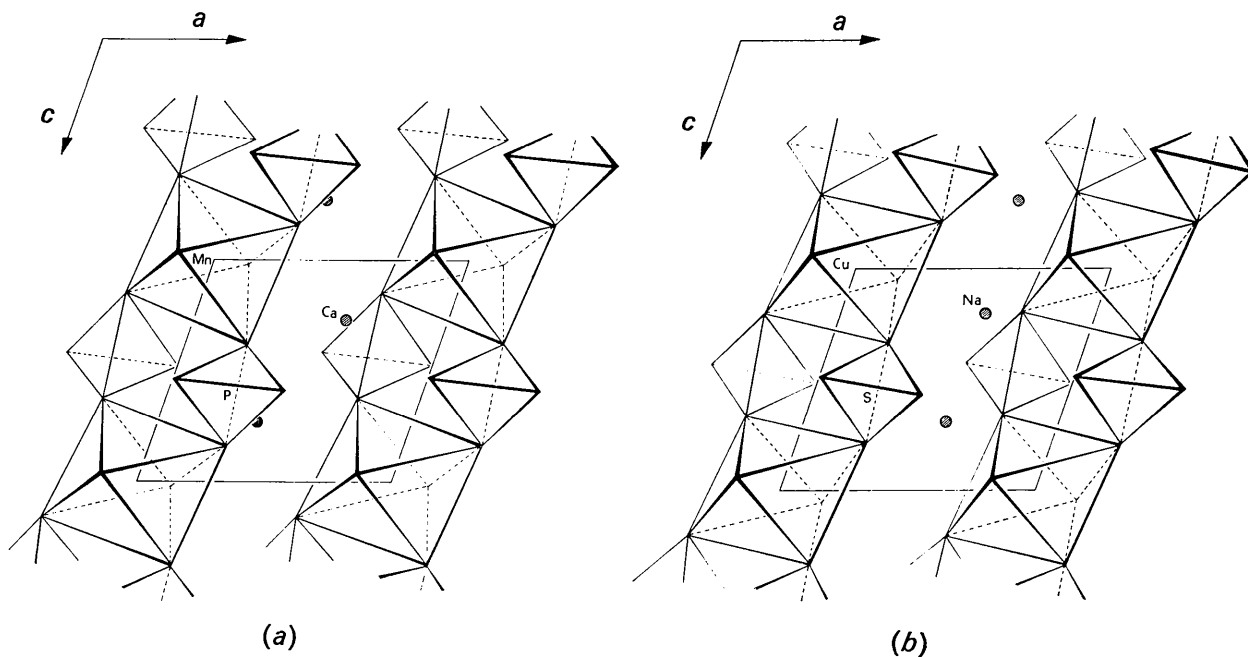


Fig. 2. A schematic view of the structures of fairfieldite (a) and kröhnkite (b) projected along the  $b^*$  axis. For kröhnkite a half cell is projected.

Table 3. Bond lengths

I	x	y	z	IV	-x	1-y	1-z
II	1+x	y	z	V	1-x	-y	1-z
III	x	y	-1+z	VI	1-x	1-y	1-z
Mn(I)-O(3) (I)				2.142 ± 0.009 Å	× 2		
-O(4) (III)				2.269 ± 0.007	× 2		
-O(W) (I)				2.180 ± 0.007	× 2		
P(I)-O(1) (I)				1.514 ± 0.007 Å			
-O(2) (I)				1.550 ± 0.007			
-O(3) (I)				1.556 ± 0.006			
-O(4) (I)				1.563 ± 0.007			
Ca(I)-O(1) (II)				2.336 ± 0.010 Å			
-O(2) (I)				2.456 ± 0.008			
O(2) (VI)				2.465 ± 0.008			
-O(3) (I)				2.475 ± 0.008			
-O(4) (III)				2.495 ± 0.011			
O(4) (V)				2.500 ± 0.007			
O(W) (II)				2.438 ± 0.007			
-O(3) (V)				2.724 ± 0.009			
O(W) (I)-O(1) (IV)				2.650 ± 0.012 Å			
-O(1) (III)				2.941 ± 0.009			

Table 4. Bond angles

O(3) (I) — Mn(I) — O(4) (III)	87° 52' ± 18'
-O(W) (I)	90 55 ± 21
O(4) (III) — Mn(I) — O(W) (I)	95 42 ± 16
O(1) (I) — P(I) — O(2) (I)	111 56 ± 20
O(3) (I)	112 38 ± 25
-O(4) (I)	109 15 ± 22
O(2) (I) — P(I) — O(3) (I)	104 36 ± 23
-O(4) (I)	110 45 ± 22
O(3) (I) — P(I) — O(4) (I)	107 29 ± 18
O(1) (II) — Ca(I) — O(2) (I)	88 59 ± 19
-O(2) (VI)	90 37 ± 20
-O(3) (I)	131 07 ± 14
-O(3) (V)	70 50 ± 17
-O(4) (III)	149 53 ± 15
-O(4) (V)	94 19 ± 21
-O(W) (I)	72 49 ± 16
O(2) (I) — Ca(I) — O(2) (VI)	72 36 ± 16
-O(3) (I)	59 47 ± 17
-O(3) (V)	86 51 ± 14
-O(4) (III)	119 41 ± 15
-O(4) (V)	140 09 ± 14
-O(W) (I)	143 11 ± 13
O(2) (VI) — Ca(I) — O(3) (I)	111 18 ± 16
-O(3) (V)	152 44 ± 13
-O(4) (III)	89 38 ± 20
-O(4) (V)	146 49 ± 11
-O(W) (I)	75 49 ± 16
O(3) (I) — Ca(I) — O(3) (V)	70 50 ± 16
-O(4) (III)	76 04 ± 17
-O(4) (V)	89 41 ± 17
-O(W) (I)	153 02 ± 13
O(3) (V) — Ca(I) — O(4) (III)	116 29 ± 16
-O(4) (V)	57 23 ± 13
-O(W) (I)	115 14 ± 15
O(4) (III) — Ca(I) — O(4) (V)	70 09 ± 19
-O(W) (I)	78 04 ± 15
O(4) (V) — Ca(I) — O(W) (I)	74 25 ± 15
O(1) (IV) — O(W) (I) — O(1) (III)	102 49 ± 19

Zanazzi, 1966) and in beraunite, 1.54 Å (Fanfani & Zanazzi, 1967).

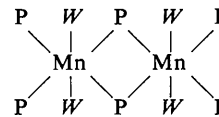
Two hydrogen bonds seem to occur between the water molecule and the oxygen atoms O(1) (IV) and O(1) (III); the distances are 2.65 and 2.94 Å, respectively; the angle between these bonds is 103°.

The charge balance according to Pauling's valence rule, assuming a completely ionic structure and taking into account the contribution of the hydrogen bonds, is as shown in Table 5.

Table 5. Charge balance of oxygen atoms

	Mn	P	Ca	-H··	··H-	Total
O(1)		$\frac{5}{4}$	$\frac{2}{7}$		$2 \times \frac{1}{4}$	$2 + \frac{3}{84}$
O(2)		$\frac{5}{4}$	$2 \times \frac{2}{7}$			$2 - \frac{15}{84}$
O(3)		$\frac{5}{4}$	$\frac{2}{7}$			$2 - \frac{11}{84}$
O(4)	$\frac{1}{3}$	$\frac{5}{4}$	$2 \times \frac{2}{7}$			$2 + \frac{13}{84}$
O(W)	$\frac{1}{3}$			$2 \times \frac{3}{4}$		$2 + \frac{10}{84}$

The structure of fairfieldite consists of chains formed by Mn octahedra and PO<sub>4</sub> tetrahedra. The resulting motif of the chain is schematically as follows (W represents a water molecule):



The chains run along the *c* axis and are connected together in the *a* and *b* directions by the coordination polyhedra around the calcium ions and by the shorter hydrogen bonds.

This structural arrangement accounts for the good {010} cleavage, but it seems not to agree with the perfect {001} cleavage which was observed by Wolfe (1940). The wide discrepancy between the measured and calculated data suggests that a new morphological study of fairfieldite, impossible to us because of the small and poorly crystallized specimen, would be useful.

#### Relationship to the kröhnkite group

The crystal structure of the triclinic minerals with formula A<sub>2</sub>B(XO<sub>4</sub>)<sub>2</sub>·2H<sub>2</sub>O can be related to the structure of the monoclinic minerals with the same general formula. The results of the present study on fairfieldite are compared with those given by Rao (1961) for kröhnkite, CuNa<sub>2</sub>(SO<sub>4</sub>)<sub>2</sub>·2H<sub>2</sub>O, this structure being generally assumed to be typical of the monoclinic family.

It is interesting to note that the unit-cell parameters are very similar in the two compounds; kröhnkite, space group *P*2<sub>1</sub>/*c*, with cell constants: *a* = 5.80, *b* = 12.67, *c* = 5.52 Å, β = 108° 36', shows nearly the same *a*, *c* and β lattice parameters as in fairfieldite, while the period along the *b* axis is nearly doubled. This feature can easily be explained by comparing the crystal structures of the two minerals: in both cases the atomic

arrangement consists of chains of coordination octahedra and tetrahedra running along the  $c$  axes with a similar configuration. Calcium ions in fairfieldite and sodium ions in kröhnkite exhibit the same role: they occupy the interstices left vacant by the chains connecting them to form a three-dimensional framework and are each surrounded by six oxygen atoms and one water molecule. From Fig. 2, where the structure of fairfieldite and a part of the atomic arrangement occurring in kröhnkite are schematically projected along the  $b^*$  axis, the resemblance of the chain motif is evident and explains the good agreement between the lattice constants in the two minerals. In fact the  $c$  translation vector represents the period in the chain and  $a \sin \beta$  the distance between two parallel chains in the  $ac$  plane.

The most remarkable difference in the two structures occurs in the reciprocal array of chains in the  $bc$  planes: in kröhnkite two parallel chains are shifted one half period along  $c$ , as is clear if one considers the presence of the glide plane at  $\frac{1}{2}b$ . This fact may account for the approximate doubling of the  $b$  parameter.

We are indebted to W. A. Henderson who kindly supplied us with the sample employed in the present investigation.

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## Crystal Structure of an Orthorhombic Modification of Methyl Stearate

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A new orthorhombic modification of methyl stearate has been found. The cell parameters are  $a = 5.61$ ,  $b = 7.35$ ,  $c = 95.15$  Å and the space group is  $Pnab$ ; the unit-cell contains 8 molecules. The carbon chains are arranged in the orthorhombic packing  $O\perp$ . The molecules form the same double sheets as have been found in the monoclinic modification. However, chains in neighbouring sheets are inclined with respect to each other.

### Introduction

During melting and freezing of methyl stearate, phenomena such as pre-melting and 'after-melting' (Smit, 1946; Simonelli & Higuchi, 1962) occur, which make the interpretation of melting curves and dilatometric measurements difficult. In order to study these phenomena, crystals of methyl stearate (purity > 99.8%) were grown from a  $CS_2$  solution at  $-15^\circ C$  to  $-12^\circ C$  and from a melt at  $37-38^\circ C$ . Two crystal forms were obtained from  $CS_2$  solution:

(a) a monoclinic form  $M_1$ , described by Aleby & von Sydow (1960) with unit-cell parameters  $a = 5.61$ ,  $b = 7.33$ ,  $c = 106.6$  Å,  $\beta = 116^\circ 47'$ ,  $d_{001} = 95.2$  Å;

(b) an orthorhombic form, which will be described in this paper.

The crystals obtained from the melt were of poor quality. A second monoclinic form  $M_2$  was found with the  $a$  and  $b$  dimensions about the same as those of the monoclinic and orthorhombic modifications from  $CS_2$ , but the  $c$  dimension was  $47.9$  Å. Some crystals from the melt appeared to be intergrowths of the orthorhombic form from the  $CS_2$  solution and a twin of the monoclinic form from the melt. The form  $M_2$  was not studied further in this investigation.

### Experimental

The unit-cell parameters and space group were determined by taking rotation and Weissenberg photographs calibrated with Al lines using  $Cu K\alpha$  radiation: